Obtaining the steady-state drawdown solutions of constant-head and constant-flux tests

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Abstract:

The solutions of constant-head and constant-flux tests are commonly used to predict the temporal or spatial drawdown distribution or to determine aquifer parameters. Theis and Thiem equations, for instance, are well-known transient and steady-state drawdown solutions, respectively, of the constant-flux test. It is known that the Theis equation is not applicable to the case where the aquifer has a finite boundary or the pumping time tends to infinity. On the other hand, the Thiem equation does not apply to the case where the aquifer boundary is infinite. However, the issue of obtaining the Thiem equation from the transient drawdown solution has not previously been addressed. In this paper, the drawdown solutions for constant-head and constant-flux tests conducted in finite or infinite confined aquifers with or without consideration of the effect of the well radius are examined comprehensively. Mathematical verification and physical interpretation of the solutions to these two tests converging or not converging to the Thiem equation are presented. The result shows that there are some finite-domain solutions for these two tests that can converge to the Thiem equation when the time becomes infinitely large. In addition, the time criteria to give a good approximation to the finite-domain solution by the infinite-domain solution and the Thiem equation are investigated and presented. Copyright © 2008 John Wiley & Sons, Ltd.

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INTRODUCTION

The Theis and Thiem equations are the well-known transient and steady-state drawdown solutions, respectively, of a pumping test conducted in confined aquifers. Both solutions are very simple and easily used to predict the aquifer drawdown distribution in practical applications. The Theis equation is derived under the conditions of an infinite extended confined aquifer, neglecting the effect of the well radius. It is important to note that the Theis equation will give an infinite drawdown as the time approaches infinity. Chen (1984) proposed a modified Theis equation for drawdown distribution in a finite confined aquifer and gave a time criterion when applying the Theis equation.

The Thiem equation can be derived either from the continuity equation coupled with Darcy's law (Todd and Mays, 2005) or the radial steady-state flow equation (Charbeneau, 2000). Many researchers have discussed the problem of steady-state flow and warn of erroneous results when applying the Thiem equation to problems of an infinite aquifer (Bear, 1979). Zaadnoordijk (1998) proposed an analytical algorithm based on the superposition principle to simulate transient flow using the Theis equation between any two given steady-state groundwater flows represented by the Thiem equation. His study

The drawdown distribution may change in response to the constant-head or the constant-flux test. The former holds a specified drawdown in the test well while the latter keeps a constant pumping rate at the pumping well. Mathematically, these two tests can be formulated as different types of boundary value problem. With or without considering the effect of the well radius, many studies have been devoted to developing analytical solutions for these problems under various boundary conditions. The objectives of this current paper are (1) to examine the drawdown solutions obtained by considering or neglecting the effect of the well radius in a finite or infinite confined aquifer for both constant-head and constant-flux tests, (2) to give physical interpretation for the convergences of the transient drawdown solutions of these two tests to the Thiem equation after a long period of aquifer test, and (3) to present criteria for the time when the finite-domain solution can be approximated by the infinite-domain solution or the Thiem equation.

DRAWDOWN SOLUTIONS OF THE CONSTANT-HEAD TEST

This section considers a constant-head test conducted in a homogeneous and isotropic confined aquifer of constant

implies that the steady-state condition is related to consideration of the effects of the well radius and the finite boundary. However, the issue of obtaining the Thiem equation from the transient drawdown solution has never previously been addressed.

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thickness. The one-dimensional radial flow equation describing the drawdown in this confined aquifer can be written as (Batu, 1998)

$$\frac{\partial^2 s}{\partial r^2} + \frac{1}{r} \frac{\partial s}{\partial r} = \frac{S}{T} \frac{\partial s}{\partial t} \tag{1}$$

where s(r, t) is the drawdown corresponding to the radial distance r from the test well and the time variable t, S is the storativity, and T is the transmissivity. The drawdown is initially assumed to be zero before the start of the aquifer test, i.e. s(r, 0) = 0.

For the constant-head test, the drawdown in the test well is maintained constant and denoted s_w . The rim of the wellbore is selected as the inner boundary and the inner boundary condition for the drawdown can be expressed as $s(r_w, t) = s_w$ where r_w is the well radius. An outer boundary condition should be provided to solve the flow equation (1). Hereafter, we will present and discuss the solution for the outer boundary specified as a zero drawdown and located at either an infinite or a finite distance from the test well. In general, the zerodrawdown boundary plays a role in supplying water if the head in adjacent area is lower than the zero drawdown. The zero drawdown is often located at the boundary of a water body, such as a river, lake, or reservoir. In addition, the drawdown solution obtained by considering or neglecting the effect of well radius will also be presented.

Infinite domain with a finite well radius

It is assumed that the outer boundary is located at infinity and the outer boundary condition is expressed as $s(\infty, t) = 0$. The solution subject to the infinite outer boundary assumption is referred to as the infinite-domain solution. Using the method of Laplace transforms, the general Laplace-domain solution of (1) can be expressed in terms of bases I_0 and K_0 , which are modified Bessel functions of the first and second kinds of order zero, respectively. The function I_0 tends to infinity under the outer boundary condition and therefore must be excluded.

Application of the inverse Laplace transforms under the inner boundary condition yields the drawdown distribution for the constant-head test as (Carslaw and Jaeger, 1959, p.335; Peng *et al.*, 2002)

$$s(r,t) = s_w \left[1 - \frac{2}{\pi} \int_0^\infty \exp\left(-\frac{T}{S}x^2t\right) \frac{Y_0(rx)J_0(r_wx) - J_0(rx)Y_0(r_wx)}{J_0^2(r_wx) + Y_0^2(r_wx)} \frac{dx}{x} \right]$$
(2)

where J_0 and Y_0 are Bessel functions of the first and second kinds of order zero, respectively, and x is a dummy variable.

The flow rate entering the test well can then be obtained by applying Darcy's law to (2). Accordingly, the wellbore flux $Q(r_w, t)$ under the condition that the

aquifer has an infinite boundary and a finite well radius is (Peng et al., 2002)

$$Q(r_w, t) = \frac{8Ts_w}{\pi} \int_0^\infty \exp\left(-\frac{T}{S}x^2t\right)$$

$$\frac{1}{J_0^2(r_w x) + Y_0^2(r_w x)} \frac{dx}{x}$$
(3)

The exponential term on the right-hand side of (2) reduces to zero as time approaches infinity; the steady-state drawdown can thus be obtained as s_w . However this result does not account for the outer boundary condition $s(\infty, t) = 0$. Such a result implies that the steady-state drawdown prevails within the entire domain and at the test well. Also, (3) shows that the wellbore flux decreases with increasing time and approaches zero for large times. This result suggests that under the infinite outer boundary condition there is not enough water to maintain a non-zero flow rate within the aquifer for a constant-head test. Thus (2) does not converge to the Thiem equation at large times.

Peng et al. (2002) proposed an improved algorithm for numerical evaluation of (2) and (3). Their results give dimensionless drawdown and dimensionless well-bore flux in tabular forms with better accuracy than Jaeger and Clarke (1942), Jacob and Lohman (1952), and Jaeger (1956). The solid line in Figure 1 is a plot of dimensionless wellbore flux (3), $Q(r_w, t)/(2\pi T s_w)$, versus dimensionless time, Tt/r_w^2S . It can be observed that the wellbore flux decreases continuously and approaches zero rather than a constant flux as time becomes infinitely large. Similar results were also discussed in Yang and Yeh (2002, Figure 3) when considering the case of single-zone formation, and Zhan and Bian (2006, Figure 5) when decreasing the leakage across an aquitard as the time becomes large.

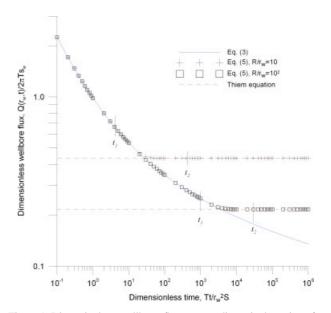


Figure 1. Dimensionless wellbore flux versus dimensionless time for the finite-domain solution (Equation (5)) with $R/r_w = 10$ and 10^2 , the infinite-domain solution (Equation (3)), and the Thiem equation. Note that both Equations (3) and (5) are solved for the constant-head test

Infinite Domain with Neglecting the Well Radius

When neglecting the effect of test well radius, that is, $r_w \to 0$, the basis K_0 of the general Laplace-domain solution tends to infinity and must be excluded. Therefore, there is no solution for this case. The transient and steady-state solutions of the drawdown for the constanthead test with and without considering the effect of well radius under the infinite domain are listed in Table I.

Finite domain with a finite well radius

In this case, a well of finite radius is considered and a finite distance R from the well is selected to represent the outer boundary where the drawdown is kept at zero. Thus, the solution under a finite outer boundary condition, denoted as s(R,t)=0, is referred to as a finite-domain solution. Carslaw and Jaeger (1959, p.332) gave the solution for this problem with Cauchy boundary conditions. By neglecting the flux components of the Cauchy boundary, a drawdown solution of the constanthead test for an aquifer with finite domain and well radius can be obtained as

$$s(r,t) = s_w \left[\frac{\ln(R/r)}{\ln(R/r_w)} - \pi \sum_{n=1}^{\infty} \exp\left(-\frac{T}{S}\alpha_n^2 t\right) \right]$$

$$\frac{J_0(r_w \alpha_n) Y_0(r \alpha_n) - Y_0(r_w \alpha_n) J_0(r \alpha_n)}{[J_0^2(r_w \alpha_n) - J_0^2(R \alpha_n)]/J_0^2(R \alpha_n)}$$
(4)

where α_n are the roots of $J_0(r_w\alpha)Y_0(R\alpha) - Y_0(r_w\alpha)$ $J_0(R\alpha) = 0$. A detailed procedure for obtaining (4) is given in Appendix A. The flow rate at wellbore can then be determined as

$$Q(r_w, t) = 2\pi T \ s_w \left[\frac{1}{\ln(R/r_w)} + 2 \sum_{n=1}^{\infty} \frac{\exp(-T\alpha_n^2 t/S)}{[J_0^2(r_w\alpha_n) - J_0^2(R\alpha_n)]/J_0^2(R\alpha_n)} \right]$$
(5)

Obviously, the steady-state wellbore flux of (5) for time approaching infinity is $Q(r_w) = 2\pi T s_w / \ln(R/r_w)$ which indeed is the Thiem equation. Figure 1 shows that the curve (5), i.e. the finite-domain solution, coincides with (3), i.e. the infinite-domain solution, at early stages of the constant-head test and asymptotically approaches the Thiem equation after a long period of time. In

other words, the infinite-domain solution can approximate the finite-domain solution when time is less than the boundary-effect time criterion t_1 , implying that the finite boundary has no effect on the wellbore flux. In addition, the finite-domain solution can be reduced to the Thiem equation when time is greater than steady-state time criterion t_2 , implying that the wellbore flux can be considered at steady state. It is worth noting that both the Thiem equation and the infinite-domain solution have the advantage of computing the drawdown solution more easily than the finite-domain solution.

Figure 1 indicates that both dimensionless time criteria Tt_1/r_w^2 S and Tt_2/r_w^2 S increase with the dimensionless distance R/r_w . The values of Tt_1/r_w^2 S and Tt_2/r_w^2 S, for instance, are 4 and 4 × 10² for $R/r_w = 10$ and 1 × 10³ and 3 × 10⁴ for $R/r_w = 10^2$, respectively, if one defines that the absolute difference in dimensionless wellbore flux between (3) and (5) is less than 10^{-5} . Thus, the time criterion for t_1 can be obtained approximately as $R^2S/10T$ based on a linear relationship between Tt_1/r_w^2 S and R/r_w for the constant-head test. Using the same approach, t_2 is $4R^2S/T$ approximately. Table II gives a list of time criteria for the finite-domain solution.

Finite domain with neglecting the well radius

In this case, the outer boundary of the radial flow system is located at a finite distance from the test well and the inner boundary of the well radius is negligible. Note that the general Laplace-domain solution for drawdown in a constant-head test contains the bases of I_0 and

Table II. The boundary-effect and steady-state time criteria for finite-domain solutions

Solution type	Boundary-effect time criterion, t_1	Steady-state time criterion, t_2	
Constant-head test	$\frac{1}{10}\tau$	4τ	
Constant-flux test	$\frac{3}{100}\tau$	8τ	
Constant-flux test when neglecting well radius	$\frac{1}{16}\tau$	$\frac{10^5}{4}\tau$	

 $\tau = R^2 S/T$

Table I. Transient and steady-state drawdown solutions of constant-head and constant flux tests

State	Constant-head test		Constant-flux test	
	Infinite domain	Finite domain	Infinite domain	Finite domain
		Considering well radius		_
Transient	Equation (2)	Equation (4)	Equation (8)	Equation (11)
Steady	*	Thiem equation	*	Thiem equation
		Neglecting well radius		
Transient	*	*	Theis equation	Equation (12)
Steady	*	*	*	Thiem equation

^{*} Indicates that the solution does not exit.

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 K_0 . The basis K_0 has to be excluded from the general solution as $r_w \to 0$ and the remaining basis I_0 cannot satisfy the boundary conditions of fixed drawdown at wellbore and zero drawdown at the finite outer boundary simultaneously. Consequently, there is no solution for this case. Table I also lists the transient and steady-state drawdown solutions for the constant-head test when the outer boundary is finite, with and without considering the effect of well radius.

Interestingly, a heat problem presented in Carslaw and Jaeger (1959, p.327) is similar to the case mentioned above but admits a solution. Analogous to the groundwater flow, their inner boundary condition has a finite drawdown instead of fixed drawdown and the outer boundary condition at a finite distance *R* from the test well has a fixed drawdown rather than zero drawdown. The solution can be written as (Carslaw and Jaeger, 1959, p.328, Equation (7))

$$s(r,t) = s_R - \frac{2s_R}{R} \sum_{n=1}^{\infty} \exp\left(-\frac{T}{S}\theta_n^2 t\right) \frac{J_0(r\theta_n)}{\theta_n J_1(R\theta_n)}$$
 (6)

where s_R represents a fixed drawdown at finite distance R, θ_n are the roots of $J_0(R\theta) = 0$ and J_1 is the Bessel function of the first kind of order one. This solution, (6), is extractable due to the fact that the inner boundary condition is finite and allowed to vary instead of being fixed in the constant-head test.

DRAWDOWN SOLUTIONS OF THE CONSTANT-FLUX TEST

This section considers a constant-flux test conducted at a confined aquifer. The initial condition for drawdown before the start of the pumping test is assumed to be zero. The drawdown solutions for (1) are presented for the cases that the outer boundary specified as zero drawdown is located at either an infinite or a finite domain and the inner boundary considers or does not consider the effect of well radius.

Infinite domain with finite well radius

The constant-flux test has a constant pumping rate Q_p . Thus, the inner boundary condition for the constant-flux test may be written as

$$2\pi r_w T \frac{\partial s}{\partial r} \bigg|_{r=r_w} = -Q_p \tag{7}$$

The drawdown solution of the constant-flux test subject to (7) and zero drawdown at infinity can be obtained by Laplace transforms as (Carslaw and Jaeger, 1959, p.338)

$$s(r,t) = \frac{Q_p}{\pi^2 r_w T} \int_0^\infty \left[1 - \exp\left(-\frac{T}{S}x^2t\right) \right] \frac{Y_0(rx)J_1(r_w x) - J_0(rx)Y_1(r_w x)}{J_1^2(r_w x) + Y_1^2(r_w x)} \frac{dx}{x^2}$$
(8)

where Y_1 is the Bessel function of the second kind of order one. Papadopulos and Cooper (1967) also

developed a drawdown solution to the constant-flux test by further considering the wellbore storage, and their inner boundary can be expressed as

$$2\pi r_w T \frac{\partial s}{\partial r}\Big|_{r=r_w} - \pi r_w^2 \frac{dH}{dt} = -Q_p \tag{9}$$

where H(t) is the well water level. Therefore, their drawdown solution can be reduced to (8) if one ignores the wellbore storage in (9). In addition, Cooper *et al.* (1967) also gave a well water level solution of the slug test by considering the wellbore storage.

Yeh *et al.* (2003) presented a closed-form solution with detailed numerical evaluations for a radial two-zone drawdown equation for groundwater under constant-flux pumping in a finite-radius well. The evaluation of (8) can be obtained if the transmissivities and storativities for the skin zone and formation zone are the same as in their article.

Figure 2 gives a graphical representation of dimensionless drawdown of (8) for dimensionless distance $r/r_w = 10^2$ and dimensionless time $Tt/(r_w^2S)$ ranging from 10^2 to 10^9 . The plot of drawdown distribution demonstrates that drawdown increases with time and becomes infinity as time approaches infinity. This implies that the outer boundary cannot provide enough water to balance continuous well pumping and the aquifer is overdrawn as time becomes infinitely large. This result indicates that the transient drawdown solution of the constant-flux test when considering the effects of well radius and the infinite domain does not reduce to the Thiem equation at very large times.

Infinite domain, neglecting well radius

The solution of the constant-flux test conducted in an infinite domain obtained by neglecting the effect of

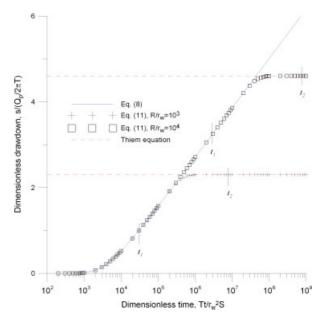


Figure 2. Dimensionless drawdown versus dimensionless time for the finite-domain solution (Equation (11)) with $R/r_w = 10^3$ and 10^4 and $r/r_w = 10^2$, the infinite-domain solution (Equation (8)), and the Thiem equation. Note that both Equations (8) and (11) are solved for the constant-flux test

well radius is analysed in this case. By applying the Laplace transforms and the asymptotic form of modified Bessel function (Abramowitz and Stegun, 1979, p.375), one obtains the drawdown equation as

$$s(r,t) = \frac{Q_p}{4\pi T} \int_u^\infty \frac{\exp(-x)}{x} dx \tag{10}$$

where $u = r^2 S/(4T \ t)$, indicating that the value of u is inversely proportional to time.

Equation (10) is the well-known Theis equation and the integral in (10) is called the well function and expressed as W(u). This function tends to infinity when time approaches infinity and/or radial distance approaches zero; thus the drawdown of the Theis equation also becomes infinite. In addition, Cooper-Jacob's equation, a special case of the Theis equation under the condition u < 0.01, should also give infinite drawdown when u approaches zero. Note that the expression for the difference in the drawdowns at any two points within the aquifer is the same as the Thiem equation if the drawdown is expressed in terms of Cooper-Jacob's equation. However, this expression is applicable only when the values of u for both points are less than 0.01. Accordingly, the locations of these two points should not be too far away from the pumping well and the expression for the drawdown difference based on Cooper–Jacob's equation does not apply to the entire aquifer domain.

The transient and steady-state drawdown solutions of a constant-flux test conducted in an aquifer with an infinite outer boundary are compared in Table I. The results for the constant-head and constant-flux tests indicate that these two tests in an infinite domain do not have steady-state solutions; thus, (2), (8), and the Theis equation cannot reduce to the Thiem equation.

Finite domain with finite well radius

Similar to the development of (4), the drawdown solution for (1) subject to (7) and the condition of zero drawdown at a finite domain can be obtained as, according to Carslaw and Jaeger (1959, p.334),

$$s(r,t) = \frac{Q_p}{2\pi T} \left[\ln \frac{R}{r} - \frac{\pi}{r_w} \sum_{n=1}^{\infty} \exp\left(-\frac{T}{S}\beta_n^2 t\right) \right]$$
$$\frac{J_1(r_w \beta_n) Y_0(r \beta_n) - Y_1(r_w \beta_n) J_0(r \beta_n)}{\beta_n [J_1^2(r_w \beta_n) - J_0^2(R \beta_n)] / J_0^2(R \beta_n)}$$
(11)

where β_n represent the roots of $J_1(r_w\beta)Y_0(R\beta) - Y_1(r_w\beta)J_0(R\beta) = 0$. It is noteworthy that the numerator of the second term in brackets of (11) equals $2/(\pi r_w\beta_n)$ when $r=r_w$ and the related drawdown solution was given in van Everdingen and Hurst (1949). The steady-state drawdown solution can easily be obtained by setting the time to infinity in (11), which is indeed the Thiem equation if the distance r is equal to the well radius.

Figure 2 shows the distribution of dimensionless draw-down of (11) versus dimensionless time for R/r_w being

equal to 10^3 and 10^4 with $r/r_w = 10^2$. It also demonstrates that the curve of (11), i.e. finite-domain solution, can be approximated by (8), i.e. infinite-domain solution, at early times and reduces to the Thiem equation after a long period of time. The former indicates the drawdown distribution that behaves as in an infinite aquifer before having the boundary effect. The latter implies that the drawdown distribution almost reaches its steady-state condition and the Thiem equation is applicable if the time is larger than the steady-state time criterion. Similar to the analysis used in the constant-head test, t_1 is obtained as $3R^2S/100T$ approximately for the constant-flux test when considering the well radius if the absolute difference of dimensionless drawdown between (11) and (8) is less than 10^{-5} . In addition, t_2 approximately equals $8R^2S/T$ if the absolute difference of dimensionless drawdown between (11) and Thiem equation is also less than 10^{-5} .

Finite domain, neglecting well radius

In this case, we consider that the outer boundary is located at some finite distance and the well radius effect is negligible. Chen (1984) gave the drawdown solution as

$$s(r,t) = \frac{Q_p}{4\pi T} [W(u) - W(U) + 2\Phi]$$
 (12)

and

$$\Phi = \sum_{n=1}^{\infty} \frac{J_0 \lfloor (u/U)^{1/2} \chi_n \rfloor}{\chi_n J_1(\chi_n)} \cdot \int_0^1 \exp \left[\frac{-U}{x} - \frac{\chi_n^2 (1-x)}{4U} \right] \frac{dx}{x}$$
(13)

where *U* is $R^2S/(4T t)$ and χ_n represents the roots of Bessel function $J_0(\chi_n) = 0$.

Both the arguments u and U in (12) and (13) are inversely proportional to time. The functions W(u) and W(U) can be expressed in terms of an infinite series and approximated as $-(0.5772 + \ln u)$ and $-(0.5772 + \ln U)$, respectively, as $u \to 0$ and $U \to 0$ after a long period of time (Todd and Mays, 2005). The difference of W(u) and W(U) reduces to $2\ln(R/r)$ where r should not be equal to zero. Moreover, the exponential term in the integrand of (13) approaches zero as $U \to 0$ and the value of Φ is negligible. As a result, the drawdown of (12) approaches the Thiem equation when time approaches infinity. The results of transient and steadystate drawdown solutions for the constant-flux test with a finite outer boundary when considering or neglecting well radius effects are also shown in Table I.

Again, Equation (12), i.e. a finite-domain solution, can be approximated by the Theis equation (10) when $U \ge 4$ (Chen, 1984) and reduces to the Thiem equation when $U \le 10^{-5}$ if the absolute difference of dimensionless drawdown is less than 10^{-5} . The time criteria t_1 and t_2 for the constant-flux test when neglecting the well radius effect are obtained as $R^2S/16T$ and $10^5R^2S/4T$, respectively, as shown in Table II.

CONCLUDING REMARKS

This paper addresses the development of steady-state solutions from transient drawdown solutions of constanthead and constant-flux tests in a finite or infinite domain, with or without considering the well radius. The results show that the finite domain condition is sufficient for the development of a steady-state solution from a transient drawdown solution. Such a condition ensures that the mass balance between extraction and supply is satisfied and the flow can reach its steady state condition within a finite domain. In addition, time criteria, which may be useful in practical applications, are provided for the approximation of the finite-domain solution by the infinite-domain solution or Thiem equation. It is found that the infinite-domain solutions can be used to determine the drawdown distribution or the aquifer parameters if the time is smaller than the boundary-effect time criterion for an aquifer with a finite domain. Similarly, the Thiem equation is applicable whenever the time is larger than the steady-state time criterion.

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APPENDIX A: DERIVATION OF EQUATION (4)

General solutions of the one-dimensional radial heat conduction equation, which is analogous to the groundwater drawdown equation, subject to Cauchy boundary conditions at the edges of a hollow cylinder were given in Carslaw and Jaeger (1959, p.332). The Cauchy boundary conditions in terms of drawdown were expressed as

$$k_1 \frac{\partial s}{\partial r} - k_2 s = k_3, \quad r = a$$
 (A1)

and

$$k_1' \frac{\partial s}{\partial r} - k_2' s = k_3', \quad r = b \tag{A2}$$

where k_1 , k_2 , k_3 , k'_1 , k'_2 , k'_3 are constant; a and b are the radial distance of the inner boundary and outer boundary in the considered region, respectively. Equations (A1) and (A2) represent the combination of the constant-head and constant-flux boundary conditions.

Using Laplace transforms, the drawdown solution based on (A1) and (A2) is

$$s(r,t) = A_1 - \pi \sum_{n=1}^{\infty} \exp\left(-\frac{T}{S}\alpha_n^2 t\right)$$

$$\frac{\lfloor k_1' \alpha_n J_1(b\alpha_n) - k_2' J_0(b\alpha_n) \rfloor C_0(r,\alpha_n)}{F(\alpha_n)} A_2(\alpha_n)$$
(A3)

where

$$\begin{split} A_1 &= \frac{-ak_3 \lfloor k_1' - bk_2' \ln(r/b) \rfloor + bk_3' [k_1 + ak_2 \ln(r/a)]}{ak_2 k_1' + bk_1 k_2' + abk_2 k_2' \ln(b/a)} \\ A_2 &= k_3 \lfloor k_1' \alpha_n J_1(b\alpha_n) - k_2' J_0(b\alpha_n) \rfloor \\ &- k_3' [k_1 \alpha_n J_1(a\alpha_n) + k_2 J_0(a\alpha_n)] \\ C_0(r, \alpha_n) &= J_0(r\alpha_n) [k_1 \alpha_n Y_1(a\alpha_n) + k_2 Y_0(a\alpha_n)] \\ &- Y_0(r\alpha_n) [k_1 \alpha_n J_1(a\alpha_n) + k_2 J_0(a\alpha_n)] \\ F(\alpha_n) &= (k_1'^2 \alpha_n^2 + k_2'^2) [k_1 \alpha_n J_1(a\alpha_n) + k_2 J_0(a\alpha_n)]^2 \\ &- (k_1^2 \alpha_n^2 + k_2^2) [k_1' \alpha_n J_1(b\alpha_n) - k_2' J_0(b\alpha_n)]^2 \end{split}$$

and α_n are the positive roots of

$$[k_{1}\alpha J_{1}(a\alpha) + k_{2}J_{0}(a\alpha)] [k'_{1}\alpha Y_{1}(b\alpha) - k'_{2}Y_{0}(b\alpha)]$$
$$-[k_{1}\alpha Y_{1}(a\alpha) + k_{2}Y_{0}(a\alpha)] [k'_{1}\alpha J_{1}(b\alpha) - k'_{2}J_{0}(b\alpha)] = 0$$

For the constant-head test with a finite well radius and outer boundary, the constants in (A1) and (A2) can be replaced by $k_1 = 0$, $k_2 = -1$, $k_3 = s_w$, $k_1' = 0$, $k_2' = 1$, $k_3' = 0$, $a = r_w$, and b = R. By careful substitution, the drawdown (4) can be obtained from (A3).